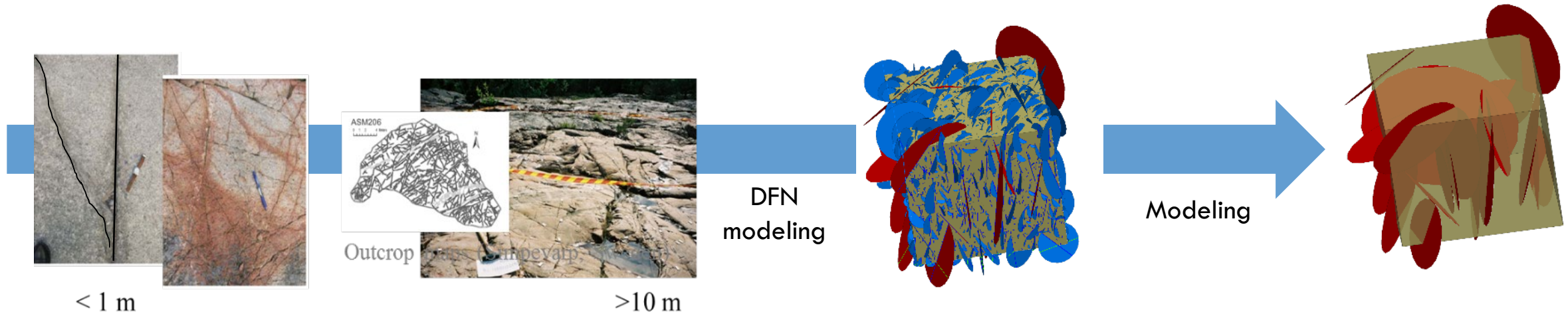


Estimating rockmass properties based on DFN methods

Caroline Darcel, Romain Le Goc, Etienne Lavoine, Diane Doolaege
Itasca Consultants, France
Philippe Davy,
Univ Rennes, CNRS, France
Diego Mas Ivars, SKB, KTH, Sweden



Objective: Introduce Discrete Fracture Network methods for rockmass modelling



Relevant complexity

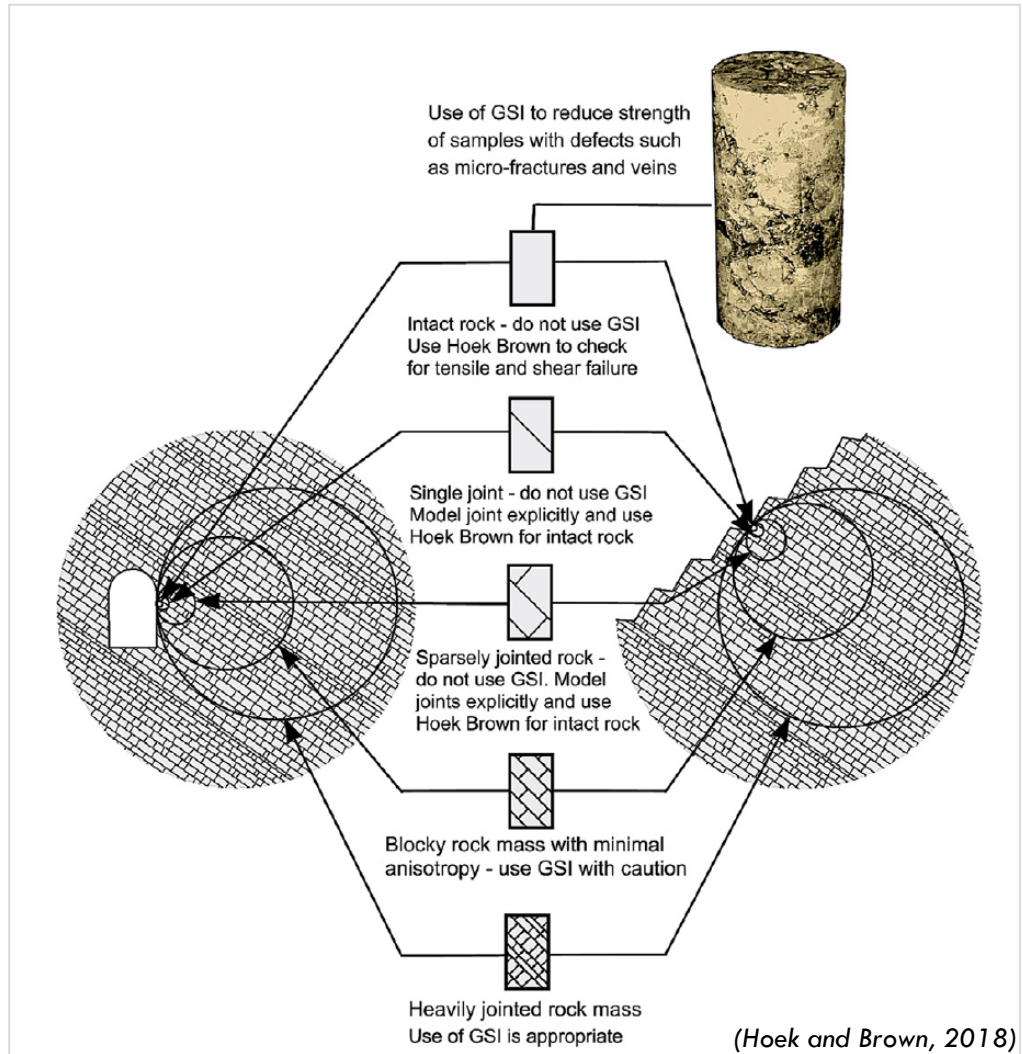
- Anisotropy
- Scale effect
- Integration from cm to km

Modeling purpose

- **Stress, strain**, hydraulic, transport, HM, thermal
- Available data
- Modeling as DEM, discrete or equivalent continuous modeling

Define which metric of a DFN is the controlling factor of rockmass elastic properties

Classical approach: fractured rock as a block assembly

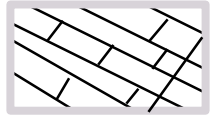


(*RQD*, *RMR*, *Q*, *GSI*) are adapted to heavily jointed rock masses

- Assembly of blocks
- Several sets of potentially infinite fractures (isotropy)
- Fractured system Scale = **spacing**
- Applicable when **model resolution** >> **spacing**

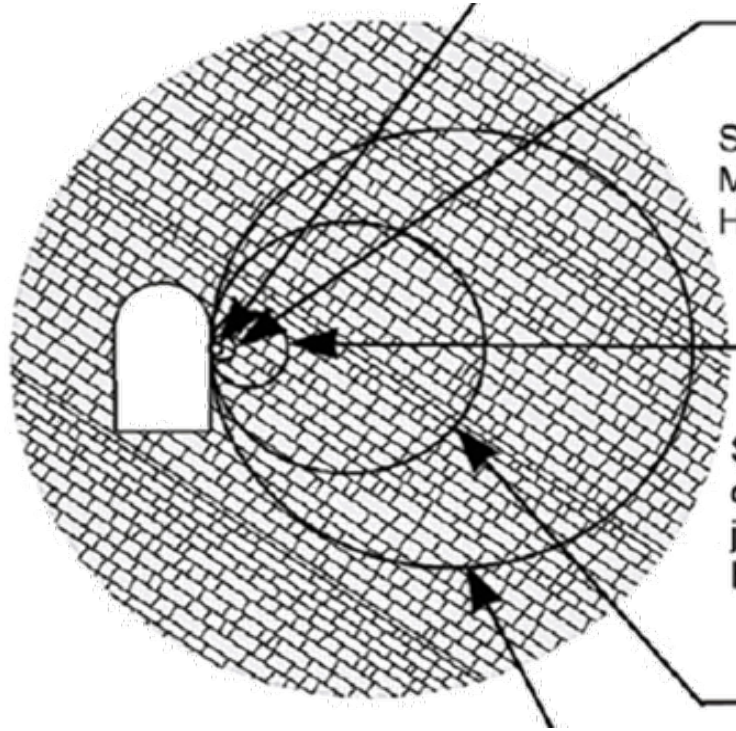
But lack of

- Density and scale (size distribution of fractures)
- 3D representation
- Anisotropy
- Quantitative description

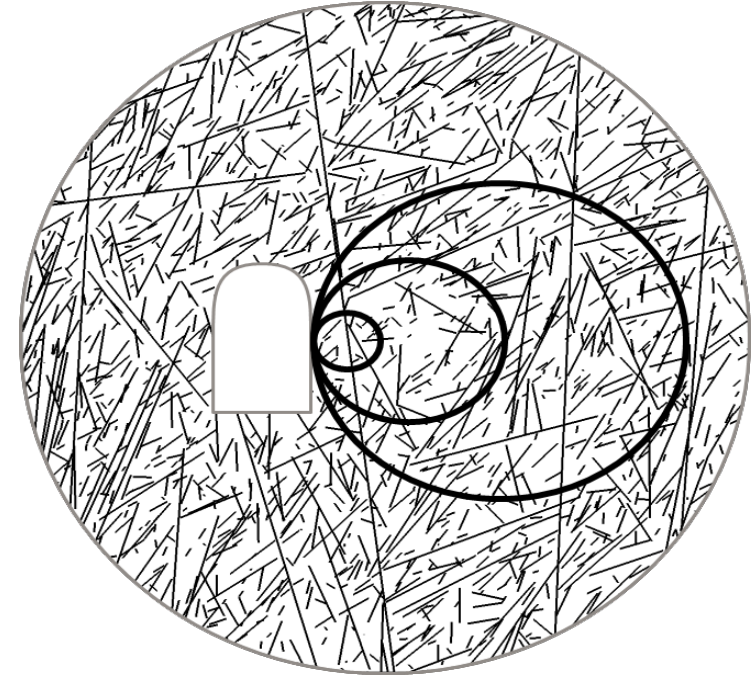


Concepts for fractured rock

Block Assembly

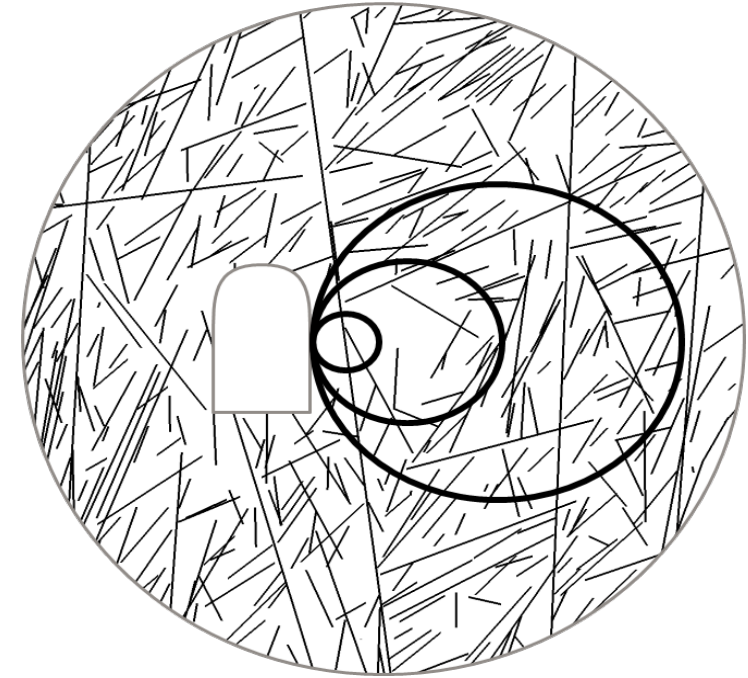
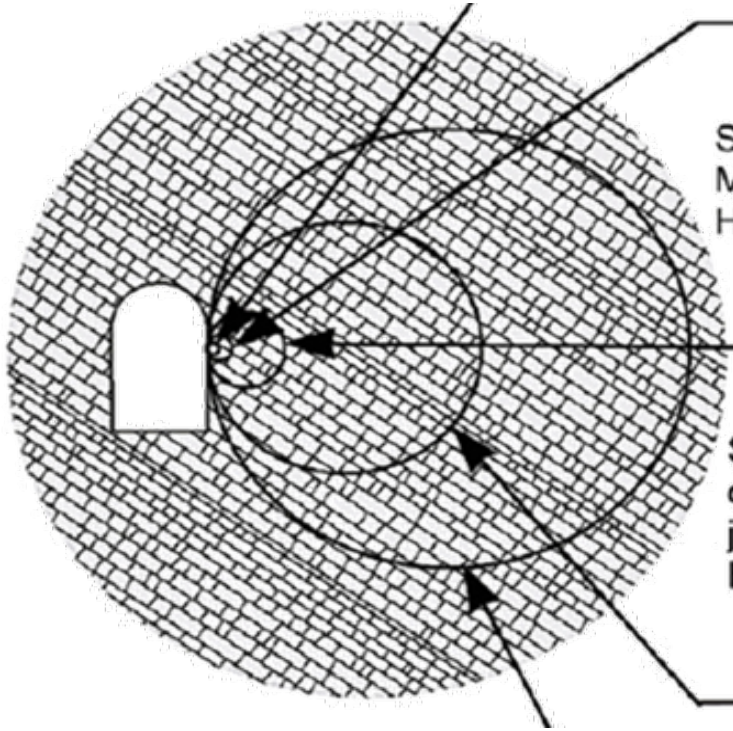


population of individual fractures



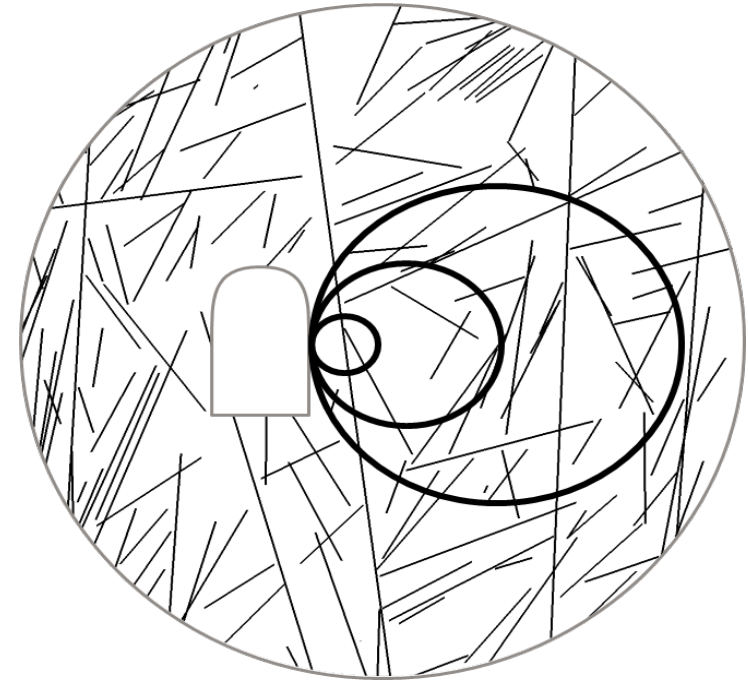
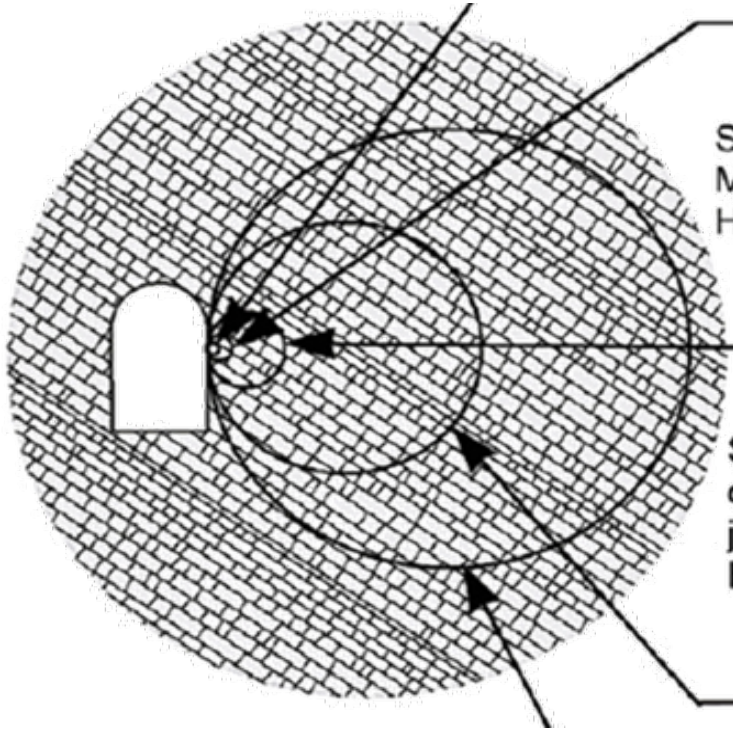
Fractured rock — block Assembly vs Network of fracture

population of individual fractures

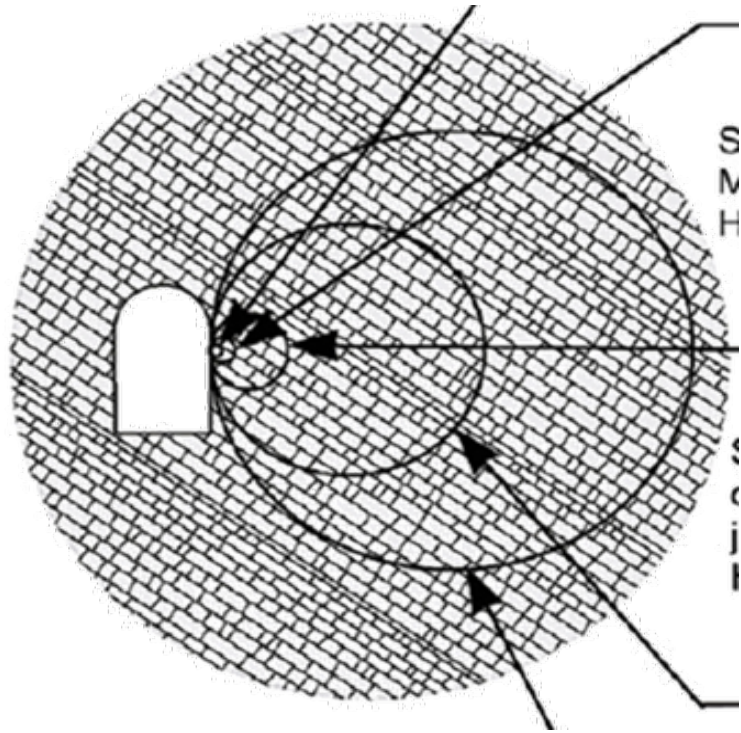


Fractured rock — block Assembly vs Network of fracture

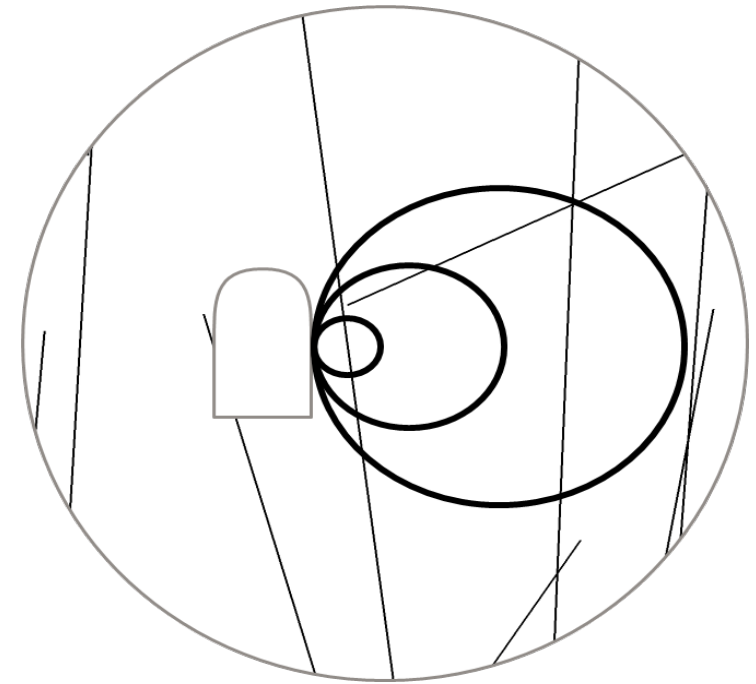
population of individual fractures



Fractured rock — block Assembly vs Network of fracture



population of individual fractures

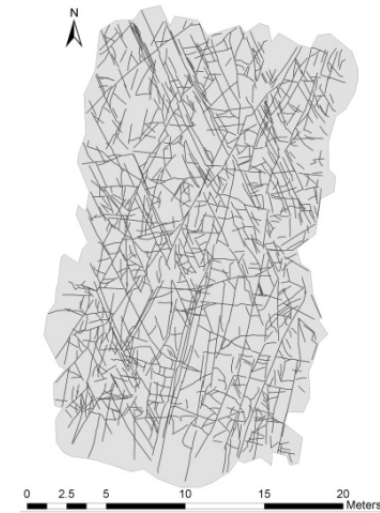


Need to assess the density vs size

Description of the fractured rock with DFN

Originally applied to

- Crystalline rocks
- Potential host for nuclear waste storage
- Connectivity, flow and transport modeling



Observation and mapping



- Mapping conditions: **Resolution** and **Censoring**
- Physical boundaries of the distribution, *min* and *max* sizes, **not directly accessible**

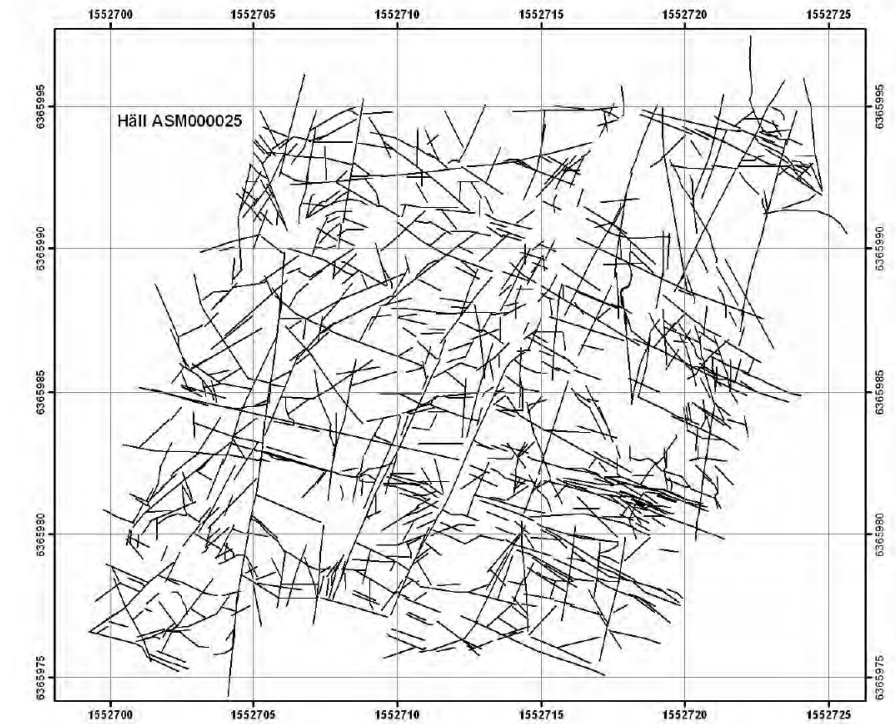
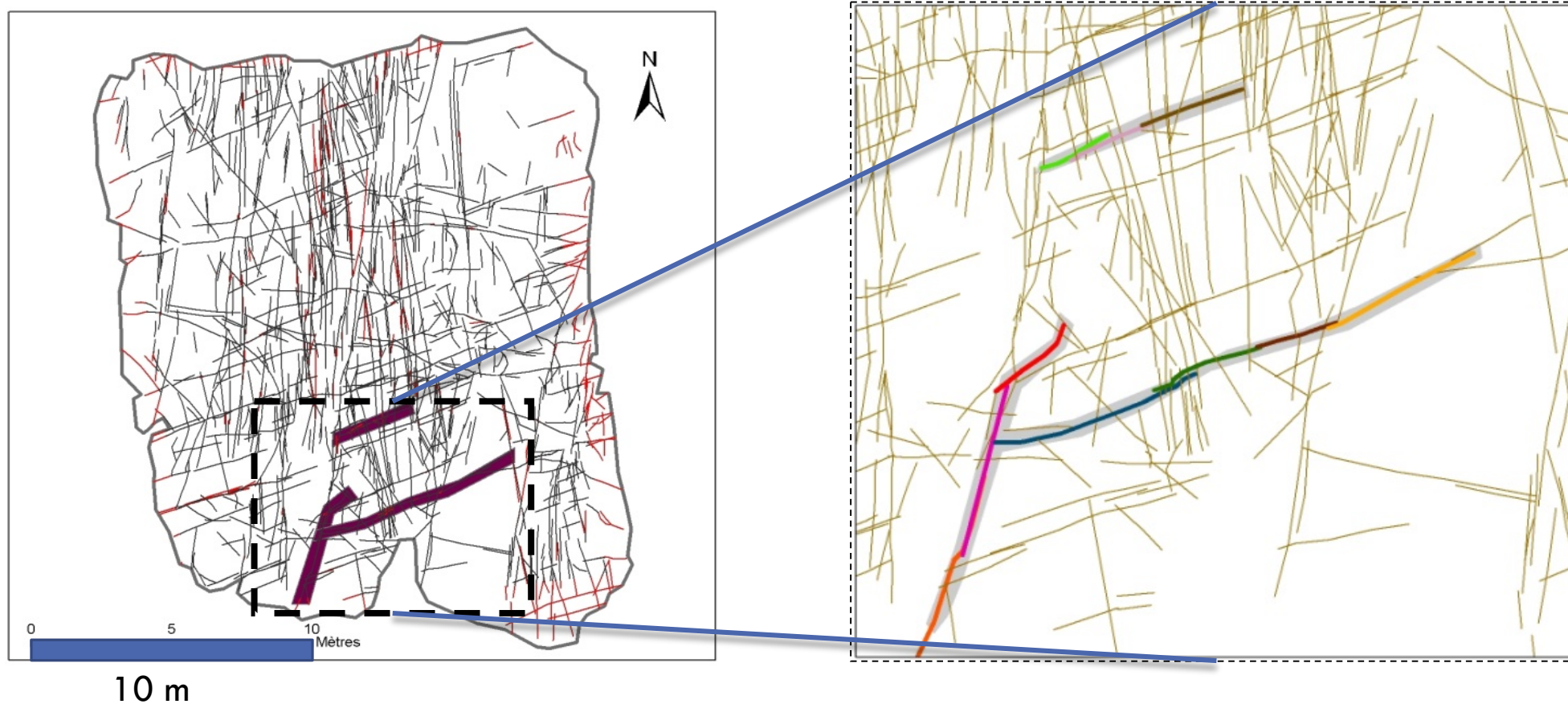


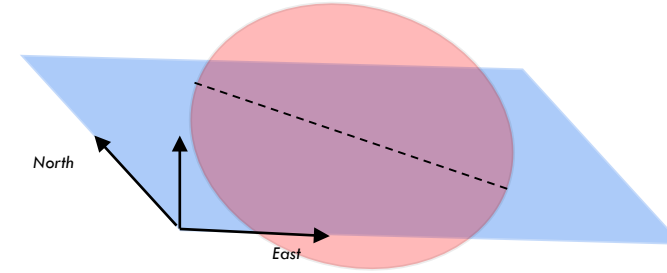
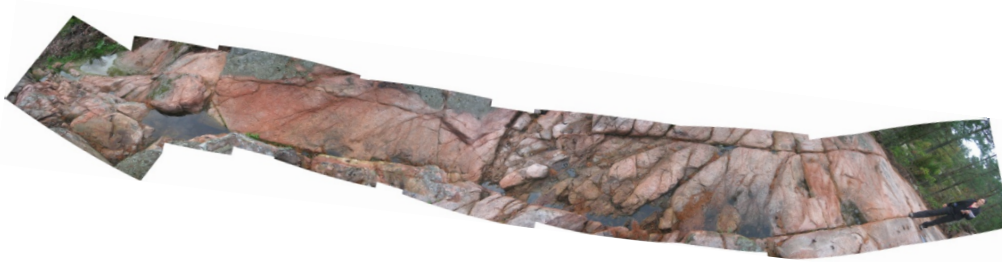
Figure 5-2. Fracture trace map of the ASM000025 outcrop.

A Fracture, what is it



A fracture in the model is an ensemble of fracture segments that define a consistent plane (mechanical coherence).

A Fracture, what is it



Roughly Planar discontinuity

Resulting from rock failure controlled by physical processes and field conditions

Can be cracks, joints, faults, shear zones, bedding planes

Lateral dimensions \gg thickness

2D planar object

Position, size, orientation, shape

Flow, transport, mechanical properties

... to the fracture population (Discrete Fracture Network)

Build the fracture size density distribution $n(l)$

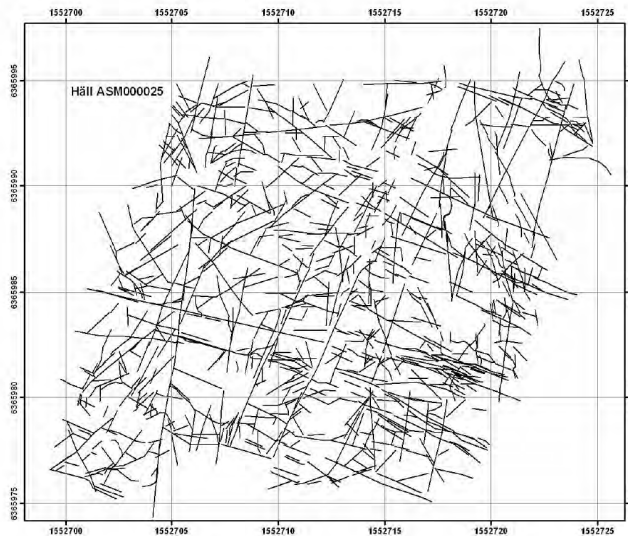
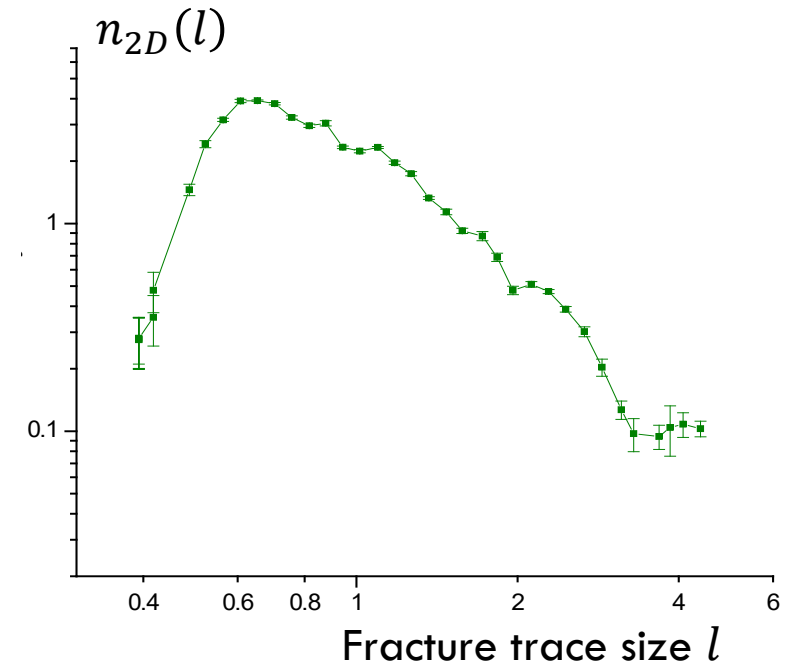


Figure 5-2. Fracture trace map of the ASM000025 outcrop.
[SKB report P-04-35]

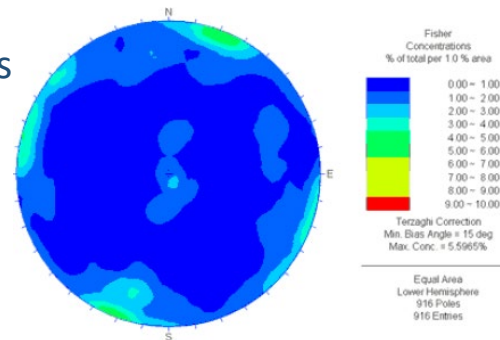
- Logarithmic binning to **count** the number of fracture traces whose size l' is in the range $l; l + dl$
- Normalized by map area and bin size

$$n_{2D}(l) = \frac{1}{area} \frac{N(l < l' < l + dl, L)}{dl}$$

- Physical boundaries of the distribution, *min* and *max* sizes, **not directly accessible**

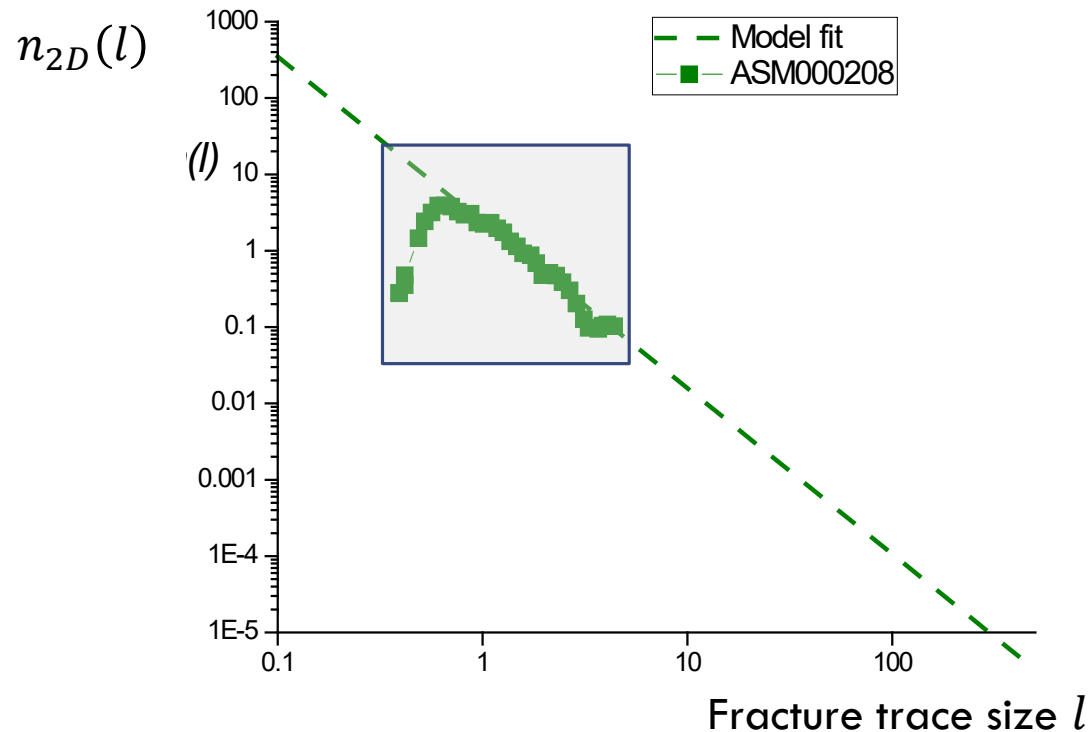


- Count orientations
- Variability ...



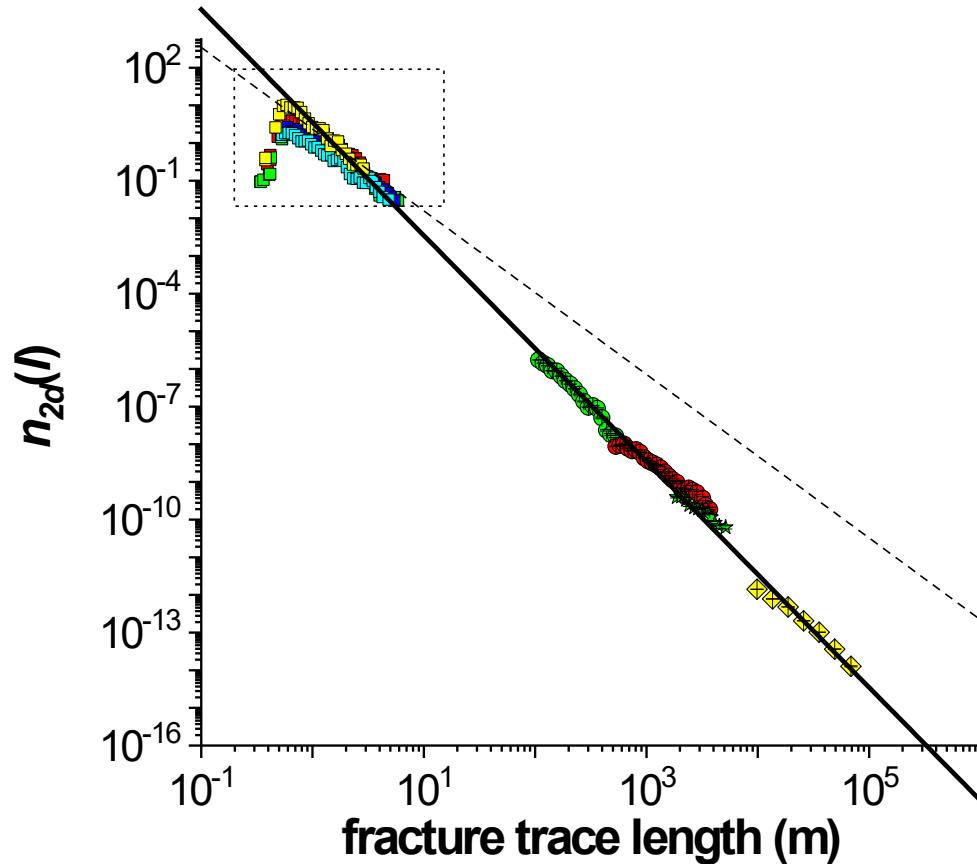
- Stereological analysis required for a 3D model

Example – 2D trace size distribution



- Power-law trend may be identified
$$n(l) = \alpha \cdot l^{-a}$$
- from a limited observation range
- parameters independent from observation range

From data to size distributions – Laxemar site

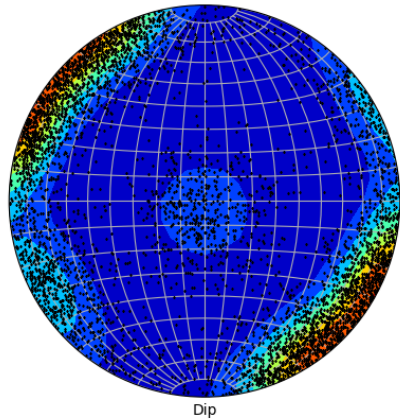
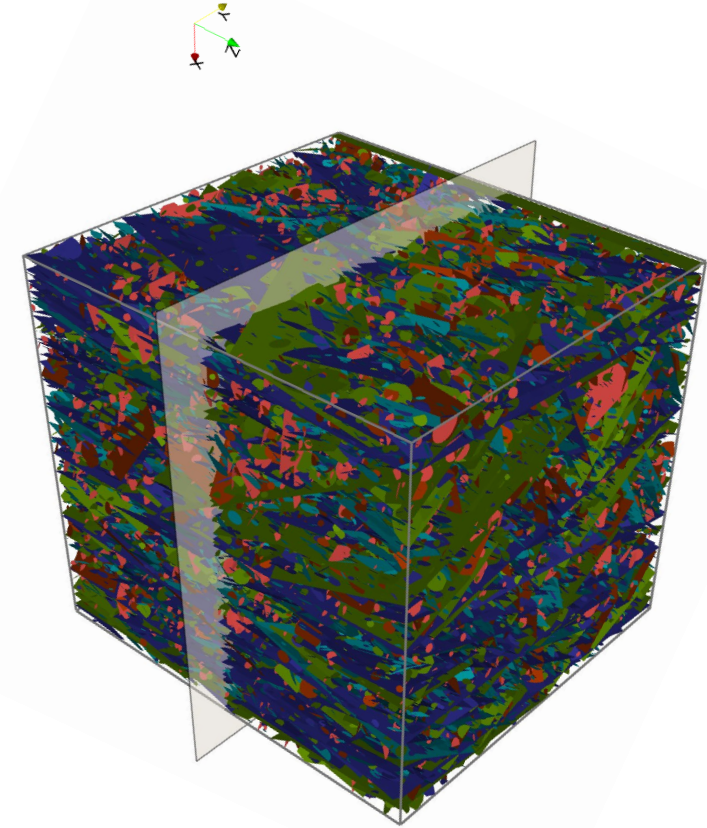
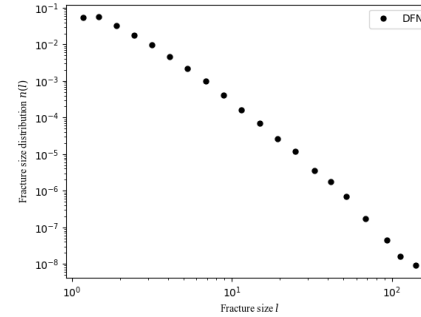
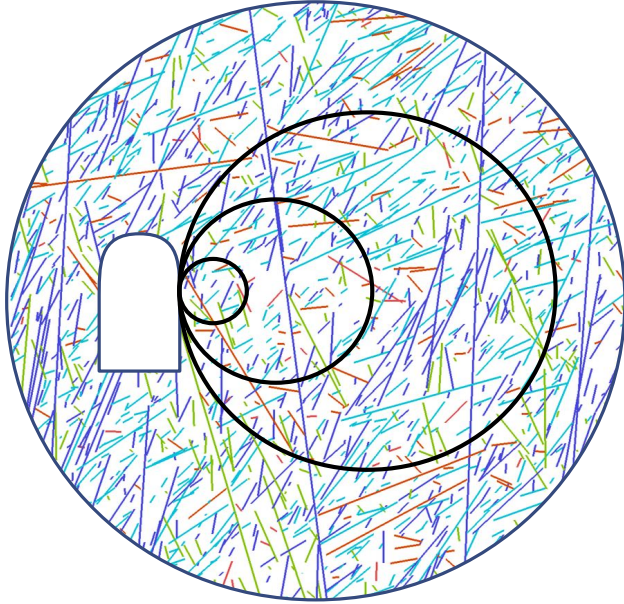


Power-law model for $n(l)$ consistent with observations

Single or double power-law model

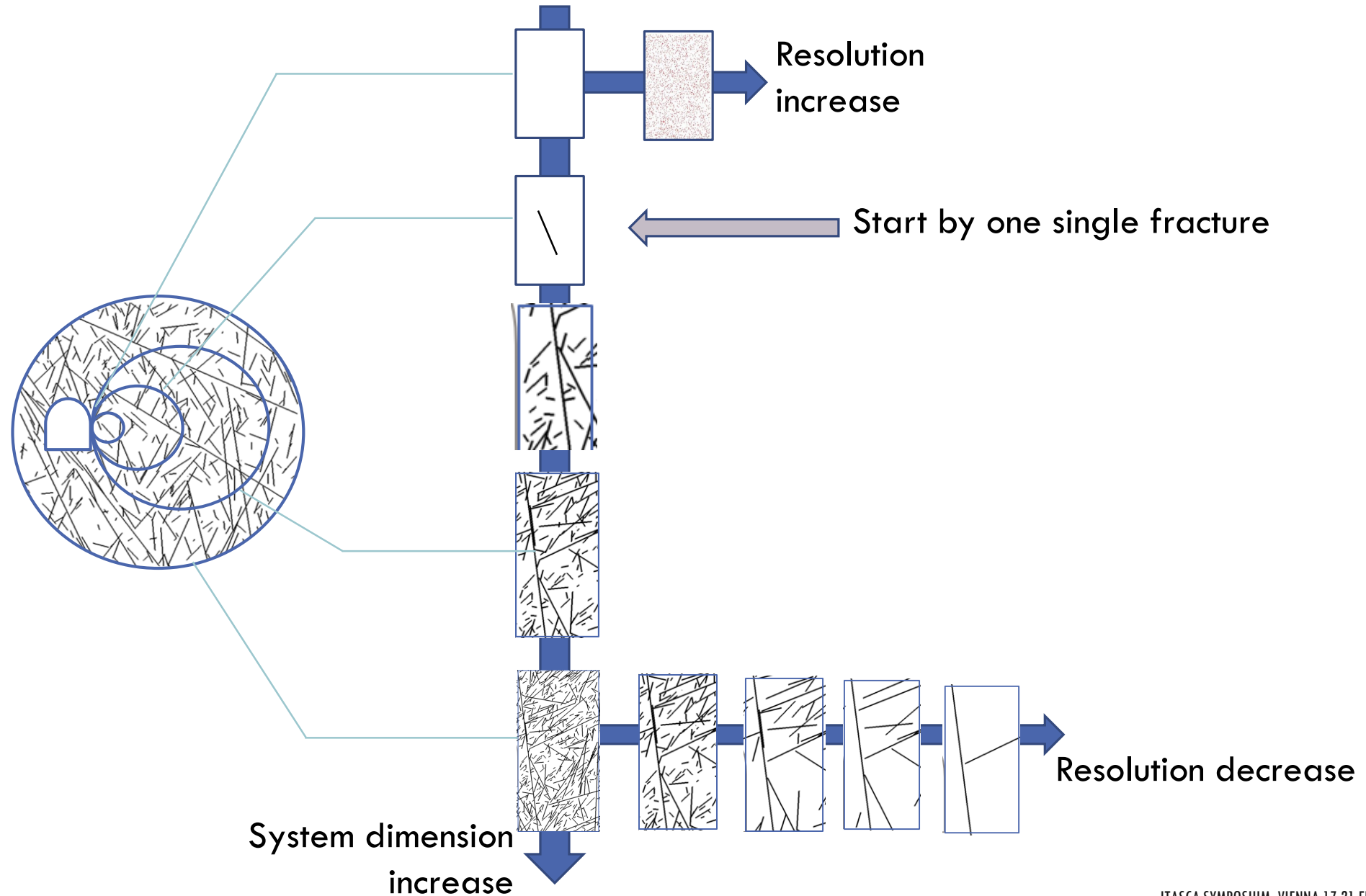
Power-law model is a good proxy to model density variation with scale

DFN in the Rockmass



- Multiscale DFN
- No a priori REV
- P_{32} (m^2/m^3) dominated by small fractures
- Connectivity related to large fractures
- Mechanical properties potentially related to size and orientations

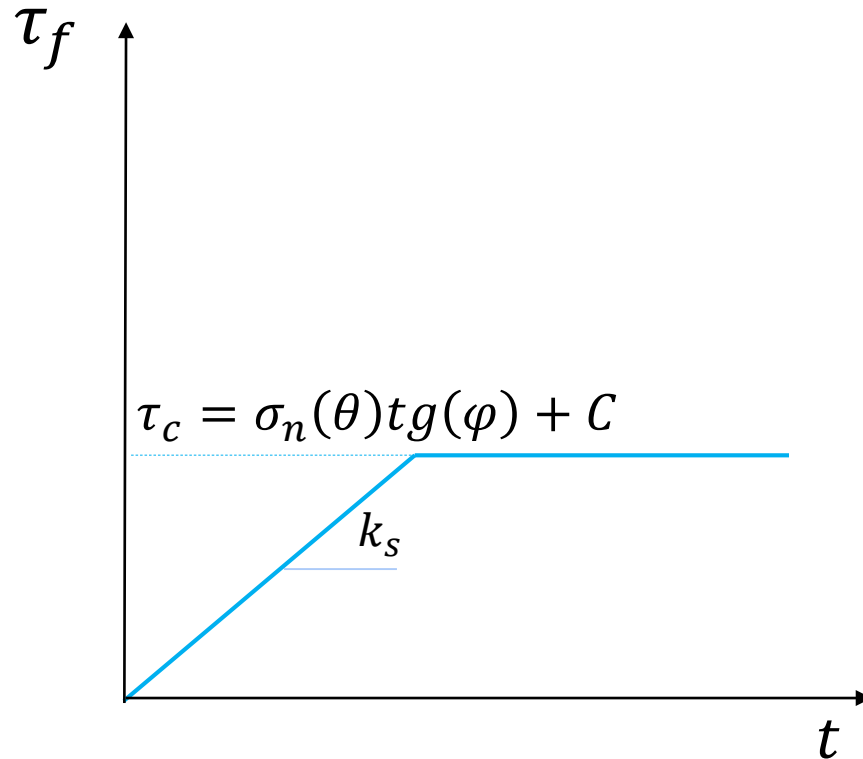
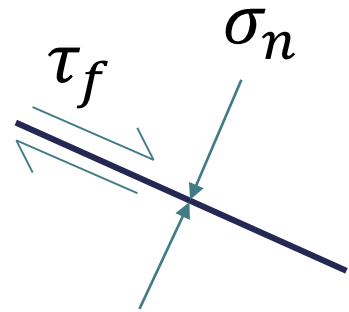
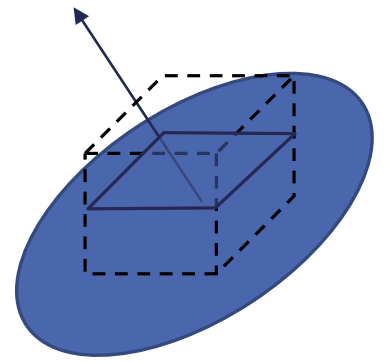
Scale, resolution, size, density



Predicting equivalent elastic properties

- From single fracture to fracture population (DFN)
- DFN: any set of disc-shaped planar fractures (multi-oriented, multi-scale)
- Elastic conditions (no damage)
- Rock matrix: isotropic elastic, Young's modulus E_m and Poisson ratio ν_m
- Fracture mechanical model
 - Coulomb slip , cohesion (c), friction (angle φ), normal (k_n) and shear stiffness (k_s)

Mechanical model - single fracture



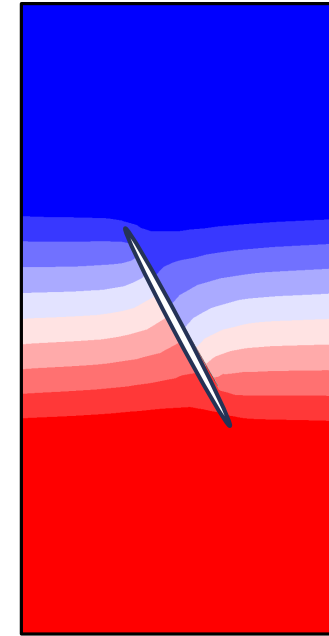
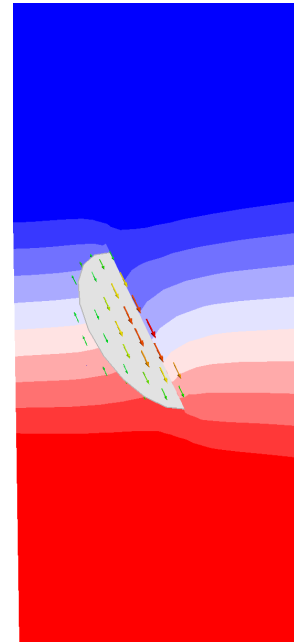
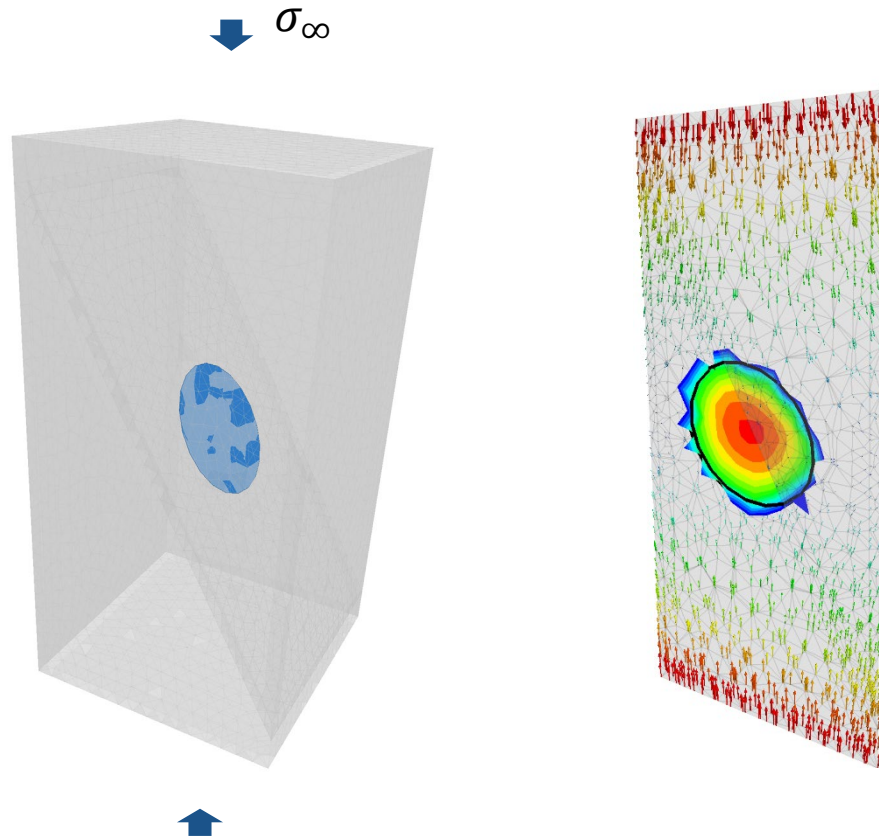
$$\tau_f = k_s \cdot t \quad \text{if } \tau_f < \tau_c$$

$$\tau_f = \tau_c \quad \text{if } \tau > \tau_c$$

$$\tau_f = 0 \quad \text{if } \varphi = 0 \text{ or } k_s = 0$$

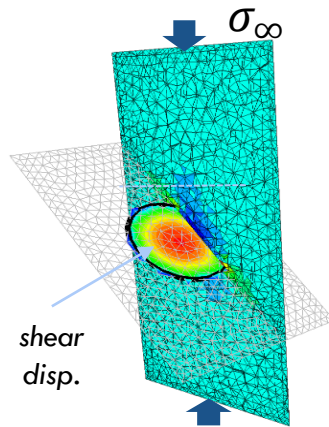
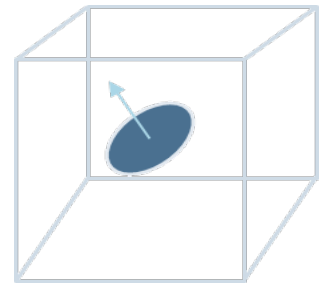
Frictionless fracture

Single fracture isolated

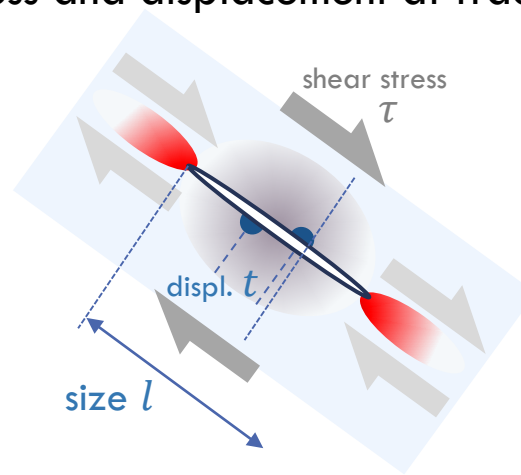


3DEC DP 5.20
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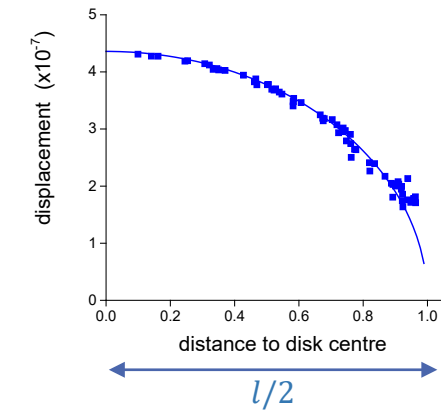
Frictionless isolated fracture



Stress and displacement at fracture



Displacement profile



Remote stress
 τ shear stress
 Intact rock
 ν_m Poisson ratio
 E_m Modulus

Fracture
 l size
 t average shear displacement

k_m equivalent matrix to fracture stiffness

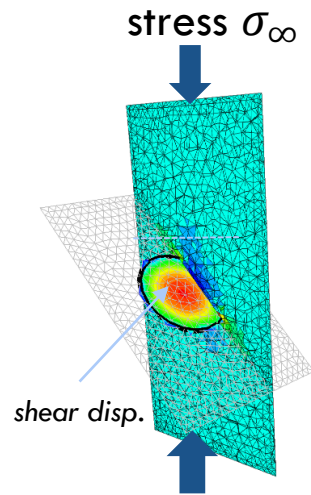
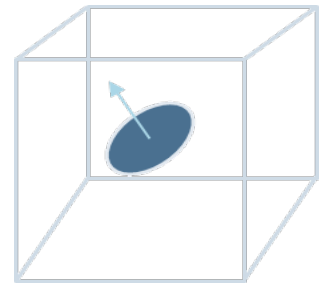
$$t = \frac{\tau}{k_m}$$

$$k_m = \frac{3\pi}{8} \cdot \frac{1-\nu_m/2}{1-\nu_m^2} \cdot \frac{E_m}{l} \sim \frac{E_m}{l}$$

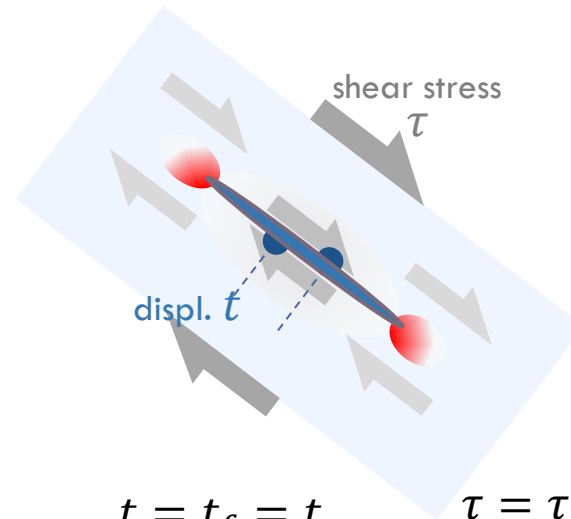
Frictional isolated fracture

Fracture friction, cohesion and stiffness terms

[Davy et al, 2018]



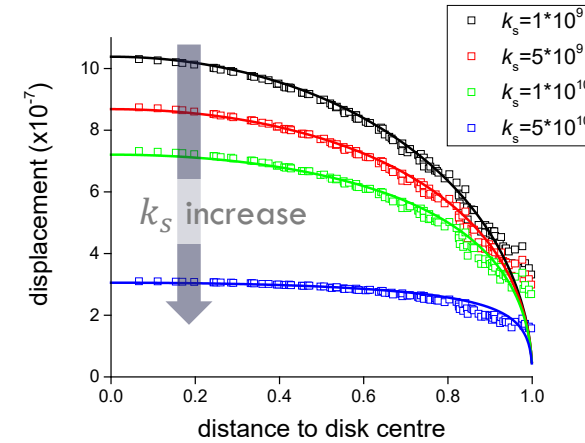
Stress and displacement at fracture



$$t = t_f = t_m \quad \tau = \tau_f + \tau_m$$

$$t = \frac{\tau}{k_m + k_s}$$

Displacement profile



Remote stress
 τ shear stress

Intact rock

ν_m Poisson ratio

E_m Modulus

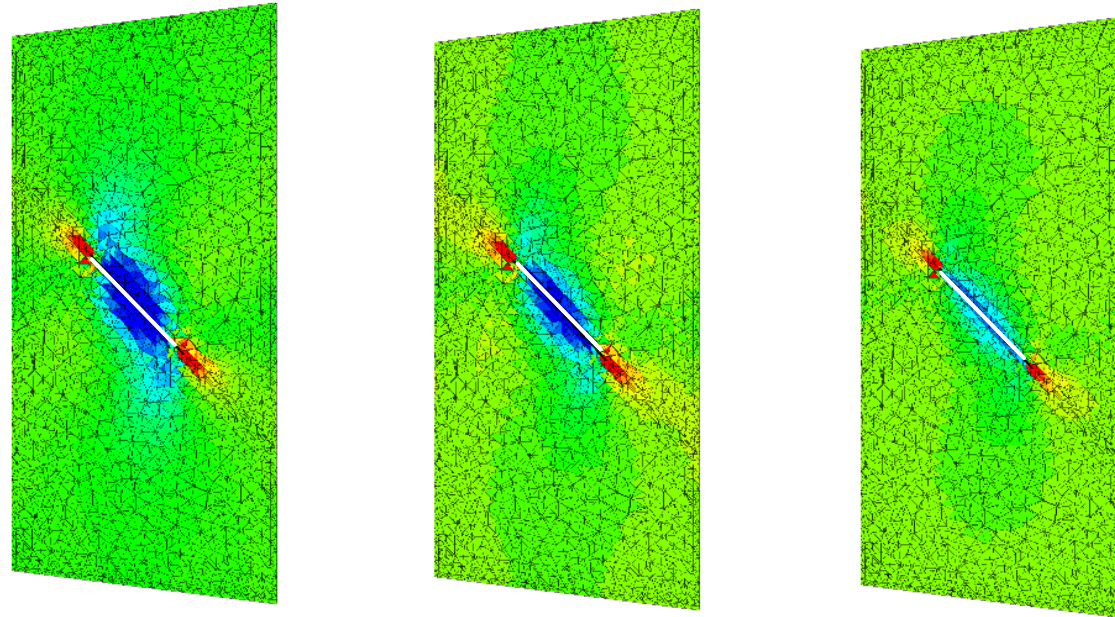
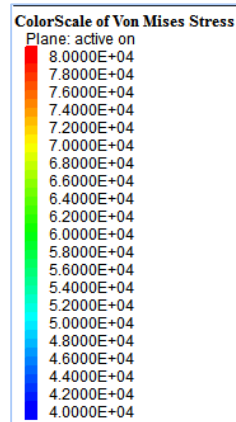
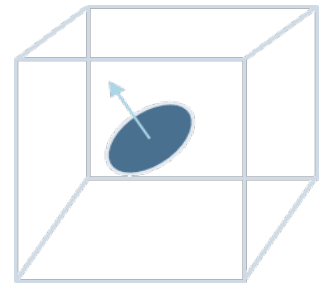
Fracture

l size

t average shear displacement

$k_m \sim \frac{E_m}{l}$ equivalent matrix to fracture stiffness

Stress perturbation around a fracture



Increasing k_s relatively to k_m
decreases the stress perturbation

Remote stress

τ shear stress

Intact rock

ν_m Poisson ratio

E_m Modulus

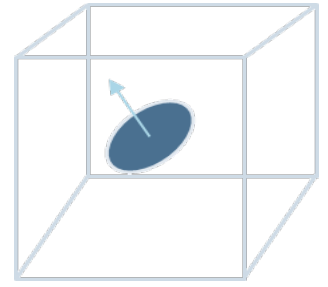
Fracture

l size

t average shear displacement

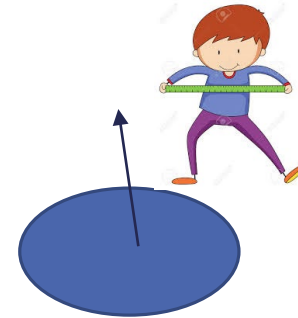
$k_m \sim \frac{E_m}{l}$ equivalent matrix to fracture stiffness

Two regimes for the shear displacement



with $l_s = \frac{E_m}{k_s}$

$$l \ll l_s \quad \rightarrow \quad t = \frac{\tau}{k_s + k_m} \approx \frac{\tau}{k_m} \propto l$$



If k_s is negligible, fracture size defines the shear displacement

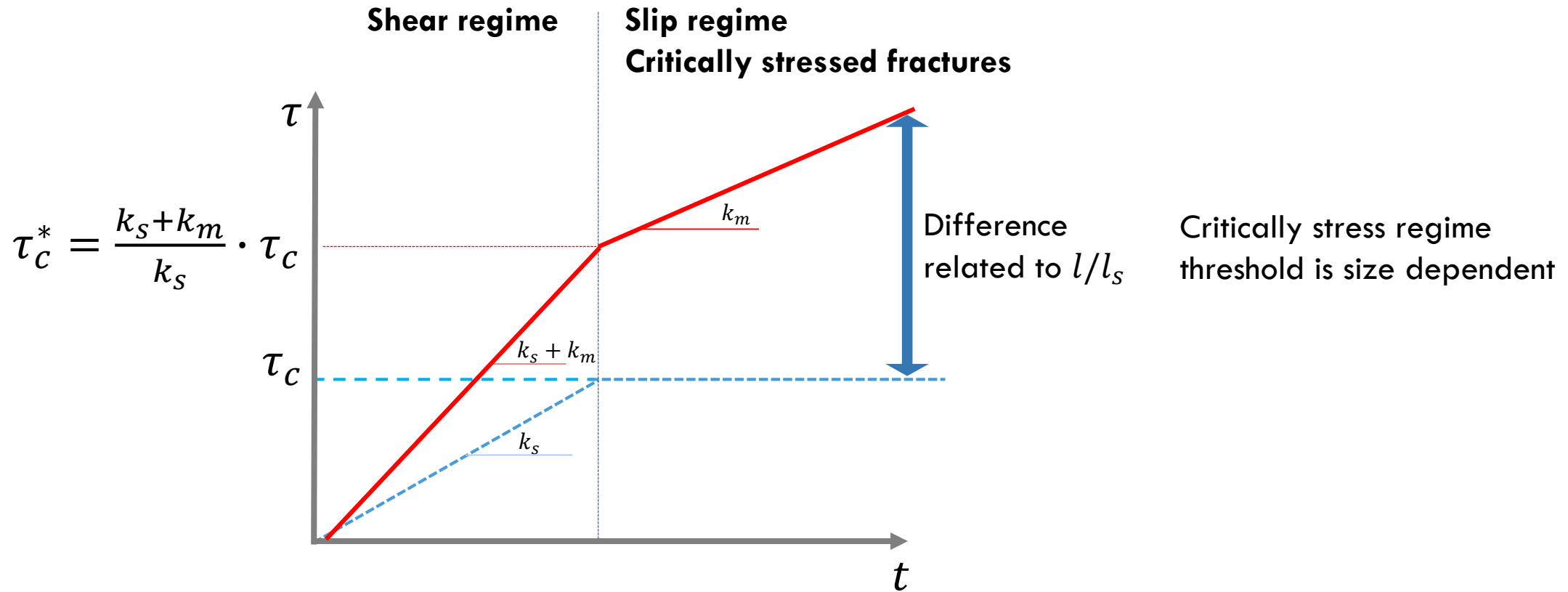
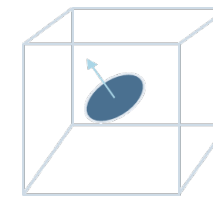
$$l \gg l_s \quad \rightarrow \quad t = \frac{\tau}{k_s + k_m} \approx \frac{\tau}{k_s}$$



If k_s is dominant, shear displacement is independent from fracture size

 Fracture sizes are critical

Shear vs slipping regime for $l \ll l_s$



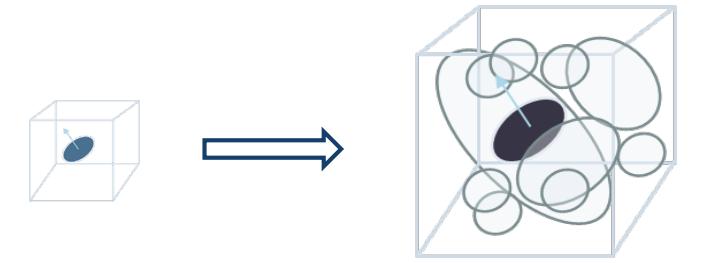
Remote stress
 τ shear stress
 Intact rock
 ν_m Poisson ratio
 E_m Modulus

Fracture
 l size
 t average shear displacement

➡ Fracture sizes are critical

$k_m \sim \frac{E_m}{l}$ equivalent matrix to fracture stiffness

From single fracture to DFN and rock mass



Rock mass with DFN

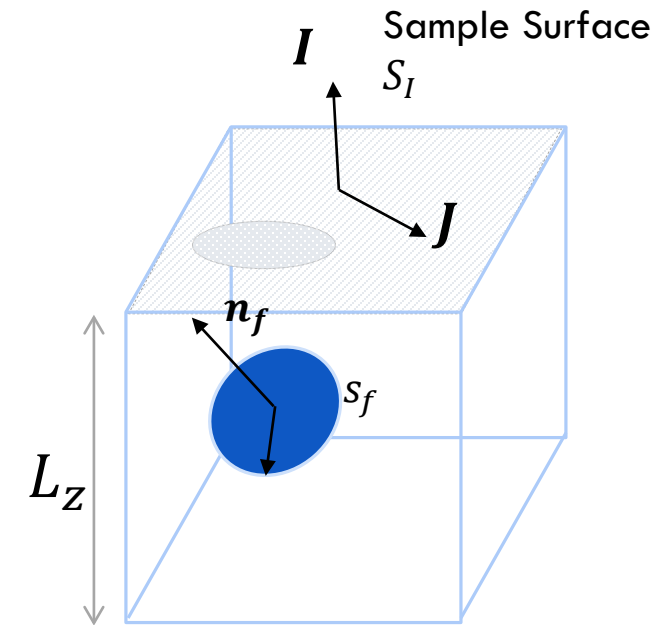
DFN contribution to rock mass strain tensor $\bar{\bar{\epsilon}}$:

Sum the contribution of each fracture f and intact rock m

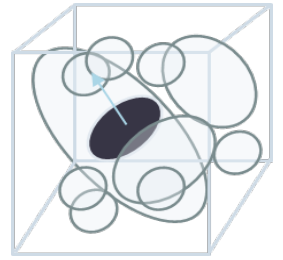
$$\epsilon_{ij} = \sum_f (\epsilon_{ij})_f + (\epsilon_{ij})_m$$

Fracture f contribution to rock mass strain

$$(\epsilon_{ij})_f = \frac{S_f(\mathbf{n}_f \cdot \mathbf{I})}{S_I} \cdot \frac{t_f(\mathbf{s}_t \cdot \mathbf{J})}{L_z} = \frac{S_f t_f}{V} \cdot (\mathbf{n} \cdot \mathbf{I})(\mathbf{s} \cdot \mathbf{J})$$



Fracture network to rock mass strain



DFN contribution to rock mass strain tensor $\bar{\bar{\epsilon}}$:

Sum the contribution of each fracture f and intact rock m

$$\epsilon_{ij} = \sum_f (\epsilon_{ij})_f + (\epsilon_{ij})_m$$

Derive effective compliance tensor components C_{ijkl} :

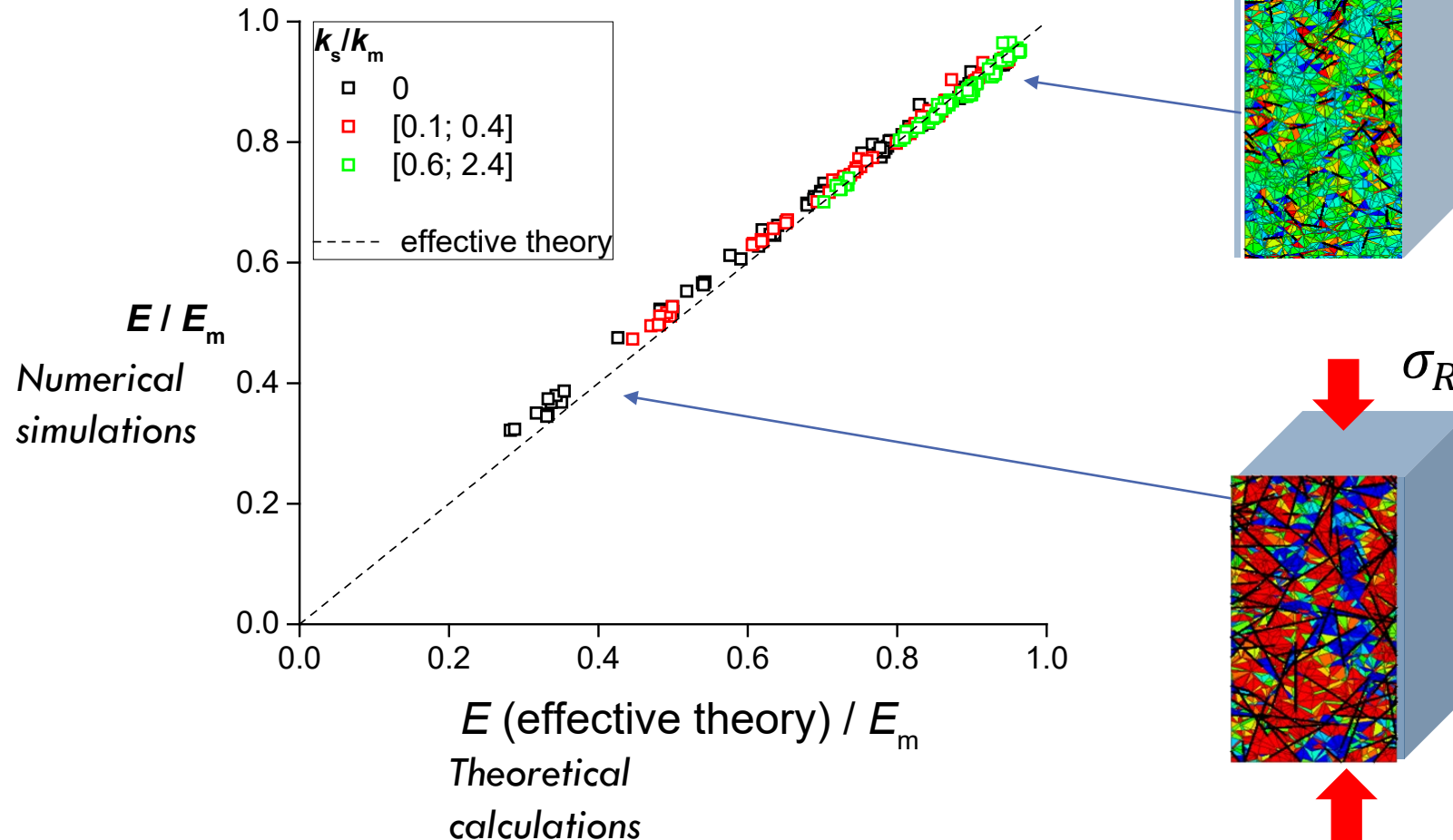
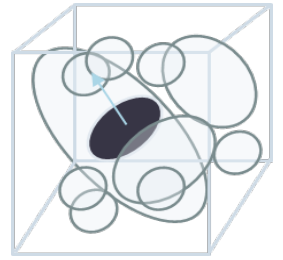
$$\epsilon_{ij} = C_{ijkl} \sigma_{kl}$$

Davy et al., 2018, Elastic properties of fractured rock masses with frictional properties and power-law fracture size distributions: JGR, v. 123, p. 6521 - 6539.

General case conditions :

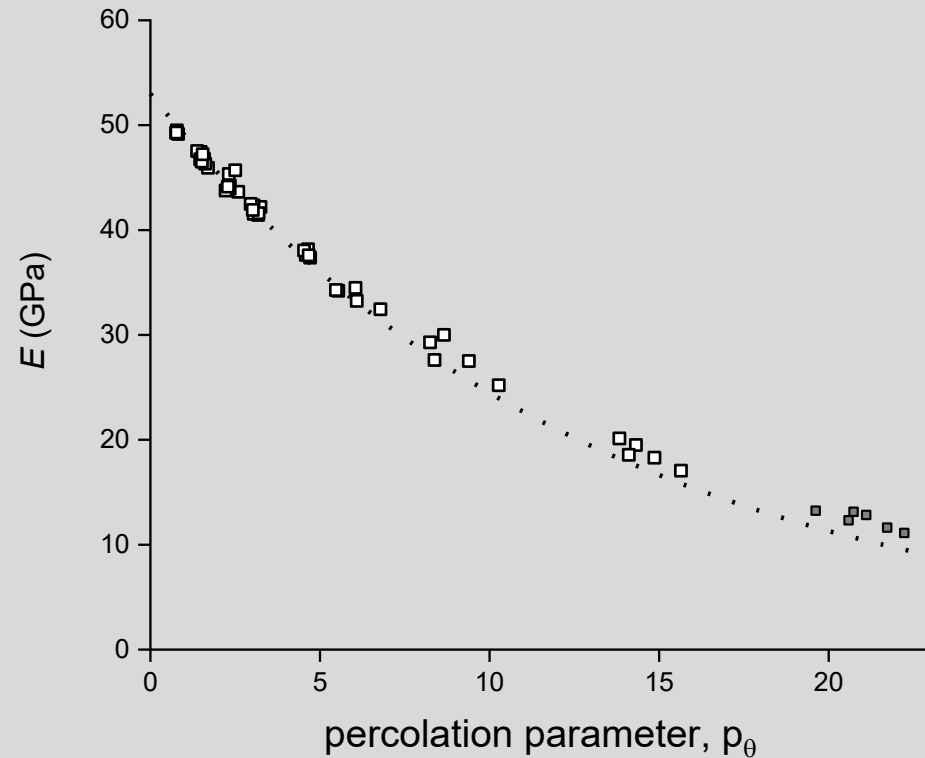
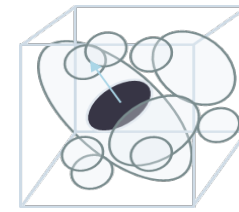
- **Shear displacement (k_s)**
- **Effective theory to account for fracture interactions for large densities**
- Change of regime for critically stressed fractures (slipping, dilation)
- Normal displacement (k_n)

Comparison to numerical simulations



Predicting E_{eff} with analytical solutions for simple cases

$k_n \gg k_s = 0$



In this case, the DFN percolation parameter $p(\theta)$ is the controlling factor of the rockmass effective elastic modulus

$$E_{eff} = E_0 \exp(-c \cdot p(\theta))$$

$$p(\theta) = \frac{1}{V} \sum_f (l_f^3 \cos^2 \theta_f \sin^2 \theta_f)$$

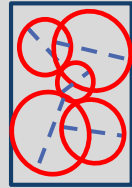
Predicting E_{eff} with analytical solutions for simple cases - $k_n \gg k_s$ constant

$$l_s = \frac{E_m}{k_s}$$

Over DFN range $l \ll l_s$

$$E_{eff} = E_m \exp(-c \cdot p(\theta))$$

$$p(\theta) = \frac{1}{V} \sum_f (l_f^3 \cos^2 \theta_f \sin^2 \theta_f)$$



p – so called percolation parameter

Over DFN range $l \gg l_s$

$$E_{eff} = \frac{k_s}{P_{32}(\theta) + k_s/E_m}$$

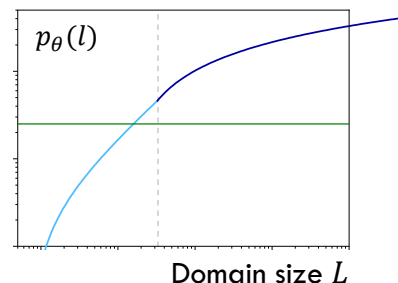
$$P_{32}(\theta) \sim \frac{1}{V} \sum_f l_f^2 \cos^2 \theta_f \sin^2 \theta_f$$



P_{32} total fracture surface per unit volume

Application to realistic multiscale DFN

Potential size effect since p is scale dependent

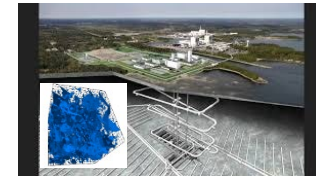


No size effect since P_{32} is scale independent

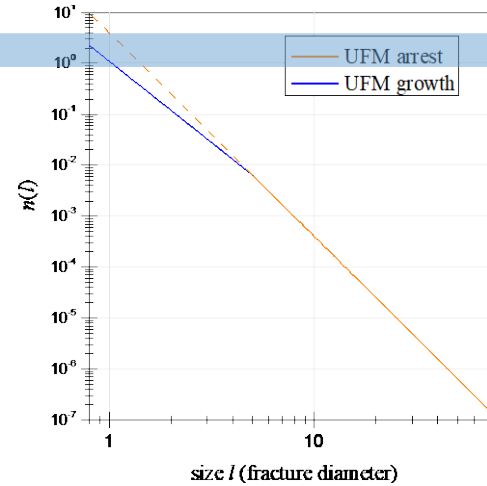
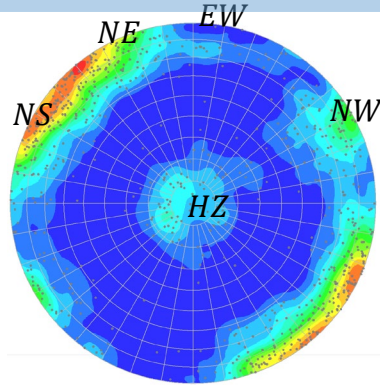
Application to site conditions — Forsmark case

- Input : generated DFN
- Input : Intact rock properties
- Input : stress state
- Output:
 - Compliance tensor
 - Scale effect
 - Level of anisotropy

Application - DFN and rock conditions SKB Forsmark site, Sweden

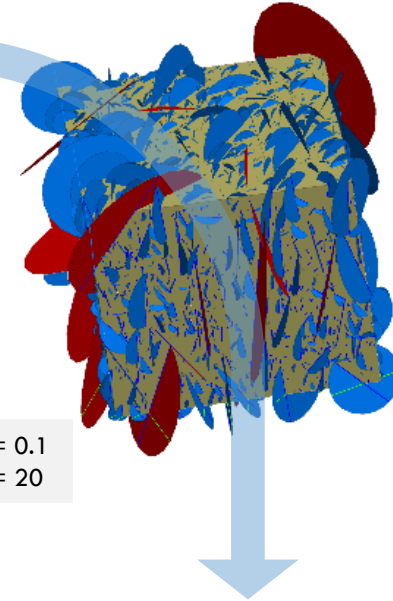


DFN (FFM01 unit)

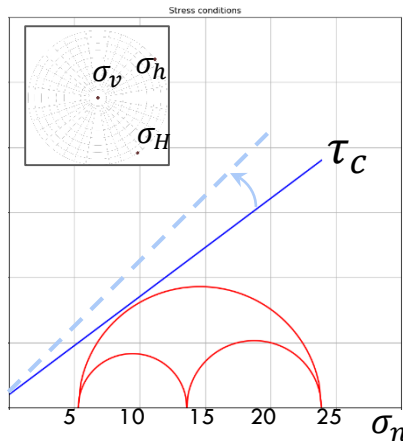


$$\frac{l_{min}}{L} = 0.1$$

$$L = 20$$



Mechanical properties



No critically stressed fractures

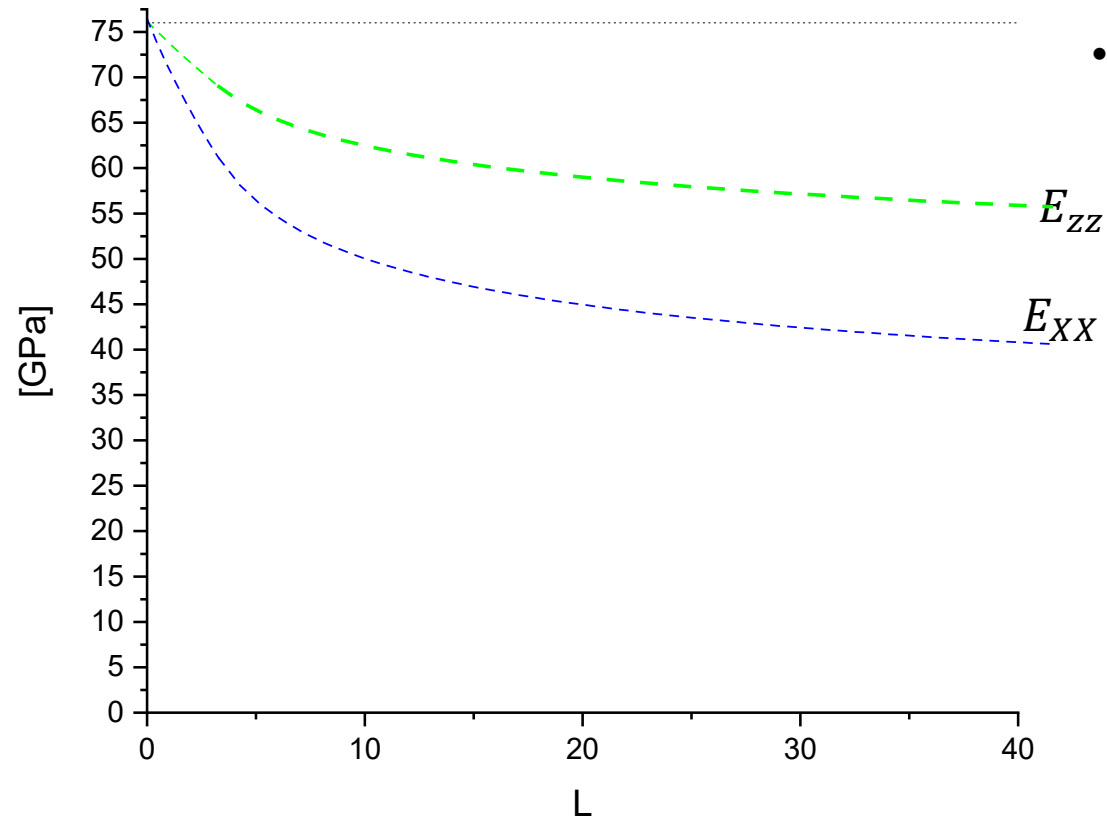
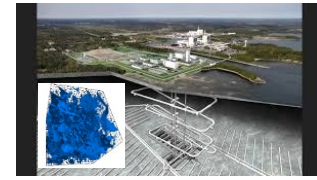
Intact Rock
 $E_m = 76 \text{ GPa}$
 $\nu_m = 0.23$

Fractures
 $k_s(\sigma_n) = 46.55 \times \sigma_n^{0.4039} \times 10^6$
 $k_n > 100k_s$

$$\begin{bmatrix} \epsilon_x \\ \epsilon_y \\ \epsilon_z \\ \epsilon_{yz} \\ \epsilon_{xz} \\ \epsilon_{xy} \end{bmatrix} = \begin{bmatrix} \frac{1}{E_{xx}} & -\frac{\nu_{yx}}{E_y} & -\frac{\nu_{zx}}{E_z} & 0 & 0 & 0 \\ -\frac{\nu_{xy}}{E_x} & \frac{1}{E_y} & -\frac{\nu_{zy}}{E_z} & 0 & 0 & 0 \\ -\frac{\nu_{xz}}{E_x} & -\frac{\nu_{yz}}{E_y} & \frac{1}{E_z} & 0 & 0 & 0 \\ 0 & 0 & 0 & \frac{1}{2G_{yz}} & 0 & 0 \\ 0 & 0 & 0 & 0 & \frac{1}{2G_{xz}} & 0 \\ 0 & 0 & 0 & 0 & 0 & \frac{1}{2G_{xy}} \end{bmatrix} \times \begin{bmatrix} \sigma_x \\ \sigma_y \\ \sigma_z \\ \tau_{yz} \\ \tau_{xz} \\ \tau_{xy} \end{bmatrix}$$

Compliance tensor $\bar{\bar{C}}$

Evolution of E_{ii} with L

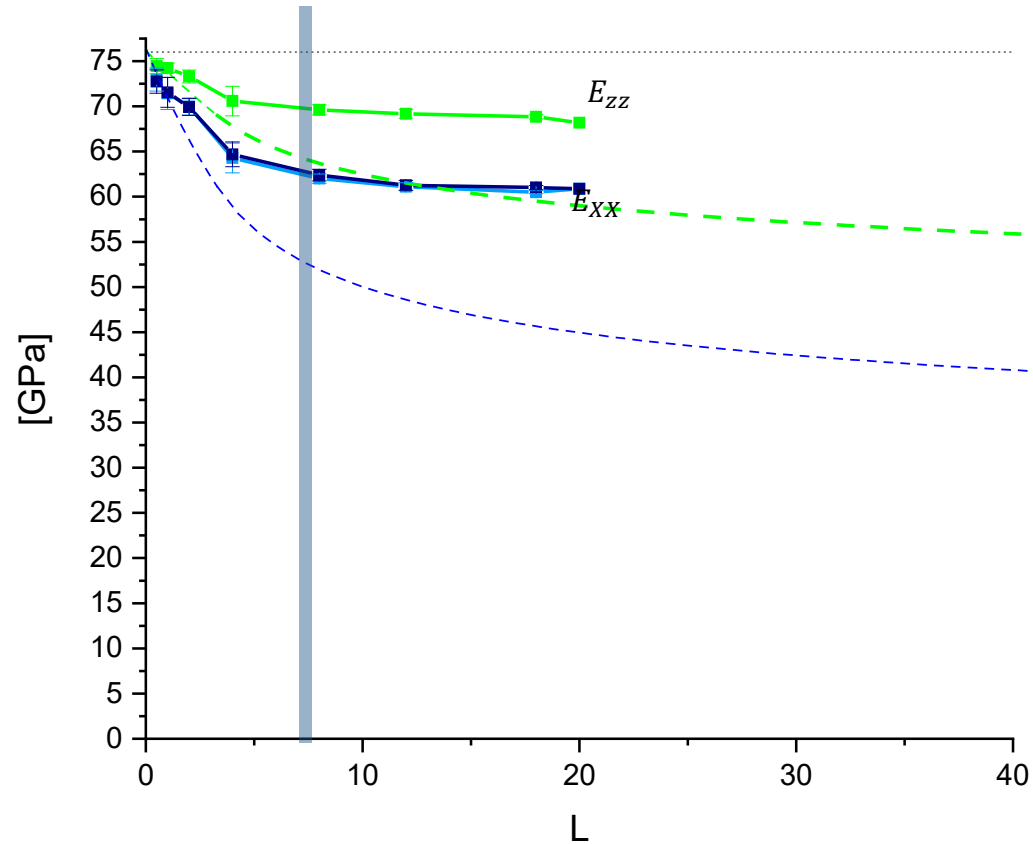
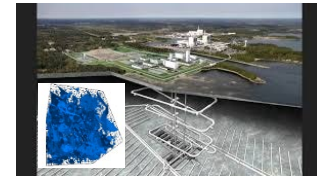


Given the DFN conditions:

- If k_s such that $l \ll l_s \rightarrow$ maximise the scaling effect

Increasing domain size L tend to put more large fractures without significantly changing P_{32}

Evolution of E_{ii} with L



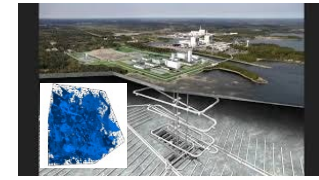
Given the DFN conditions:

- If k_s such that $l \ll l_m \rightarrow$ maximise the scaling effect

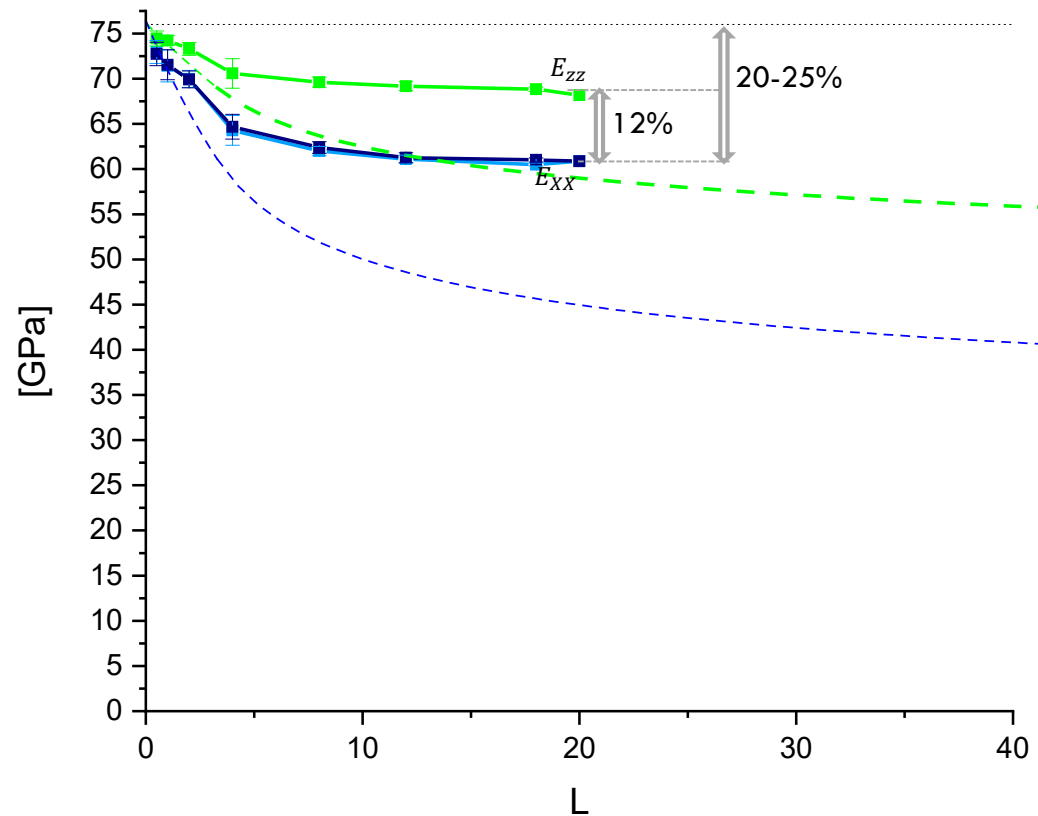
With current mechanical properties

- $\langle k_s \rangle = 3.4e10 \text{ GPa} \cdot \text{m}^{-1}$
- $1.5 \text{ m} \leq l_s \leq 3.5 \text{ m}$
- Decrease of E_{ii} with L up to $\sim 10 \text{ m}$.
- E_{xx} decrease from 76 GPa to about 62 GPa, i.e. about 25%. *

Evolution of E_{ii} - Anisotropy

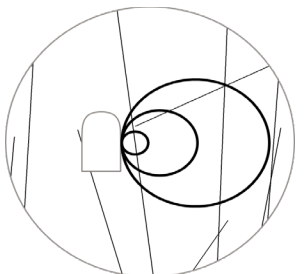
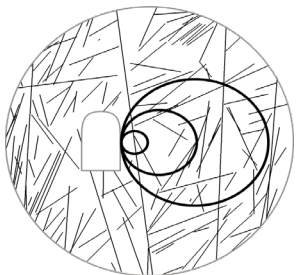
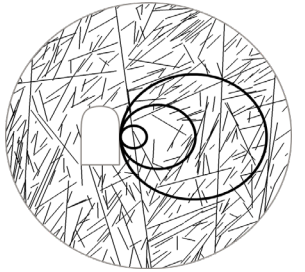
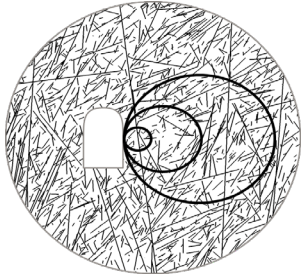


1 DFN	
l_{min}	= 0.1
L	= 20
Ori	= FFM01
k_s	(σ_n)
k_n	$\approx \infty$



- E_{ii} variations : 60 to 70 Gpa (about 12%)
- E_{zz} less affected by fractures than horizontal E_{hz}
- (Horizontal directional E_{hz} consistent with fracture sets NE and NW, less affected by fracture shearing are at trend 45°)

SUMMARY



- DFN representation of rockmass help to integrate multiscale fracture distribution
- Rockmass effective properties can be derived and controlling factors – as a combination between mechanical, geometrical and scale - identified
- Extent of scale effect and anisotropy can be quantified
- Tool (DFN.lab) to integrate DFN in geomechanical models

